



WROCLAW UNIVERSITY OF ENVIRONMENTAL AND LIFE SCIENCES

**COMPARISON OF THE ACCURACY OF TWO HIGH
RESOLUTION GLOBAL GEOPOTENTIAL MODELS:
EGM08 AND EIGEN-6C4.
CASE STUDY AT THE AREA OF POLAND**

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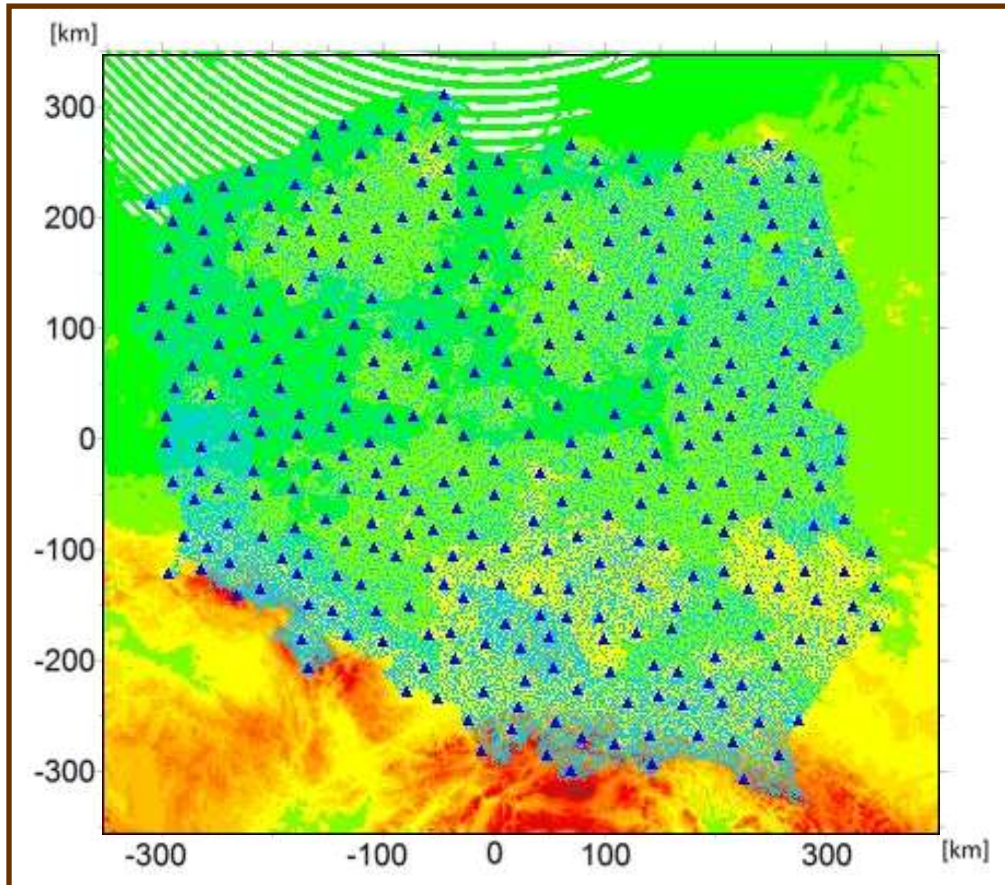
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THE ANALYSED GLOBAL GRAVITY MODELS

Model name	Year of development	Degree Order	Semi-major axis a GM	Used data
EGM2008	2008	2190 2159	6378137.00 m $3.986004418 \times 10^{14} \frac{\text{m}^3}{\text{s}^2}$	ITG-GRACE03S (180/180) 5' × 5' free-air gravity anomalies grid (global grid formed from terrestrial, altimetry-derived and airborne gravity data)
EIGEN-6C4	2014	2190 2190	6378136.46 m $3.986004415 \times 10^{14} \frac{\text{m}^3}{\text{s}^2}$	LAGEOS (30) GRACE (130) GOCE SGG Txx + Tyy + Tzz + Txz (235) 2' × 2' free-air gravity anomalies grid (altimetry over the oceans and EGM2008 over continents)

On the continents for wavelengths beyond spherical harmonics degree 235,
EIGEN-6C4 is basically a reconstruction of EGM2008

USED TEST DATA



 33330 Gravity points referred to the International Gravity Standardization Network 1971 (IGSN71)

HIGH ACCURACY GNSS/LEVELLING POINTS
(estimated error of height anomaly $\pm 1cm$)

- ▲ Selected points of networks EUREF-POL, POLREF, EUVN, ASG-EUPOS-EA - Adjustment 2011, (241 points provided by CODGiK)
short name of the group: **ASG 2011**

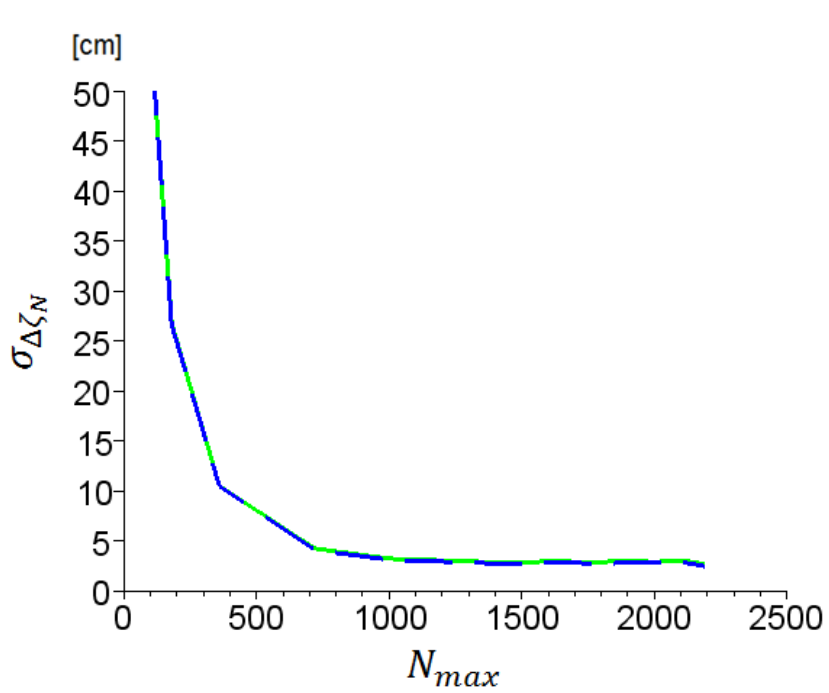
BASIC ACCURACY PARAMETERS OF THE ANALYSED MODELS

	EGM2008	EIGEN-6C4
max($\Delta\zeta$) [cm]	9.8	4.1
min($\Delta\zeta$) [cm]	-17.4	-17.8
mean($\Delta\zeta$) [cm]	-3.9	-4.5
$\sigma_{\Delta\zeta}$ [cm]	2.7	2.5
max($\Delta\delta g$) [mGal]	64.4	63.8
min($\Delta\delta g$) [mGal]	-20.4	-20.8
mean($\Delta\delta g$) [mGal]	13.4	13.1
$\sigma_{\Delta\delta g}$ [mGal]	3.7	3.7

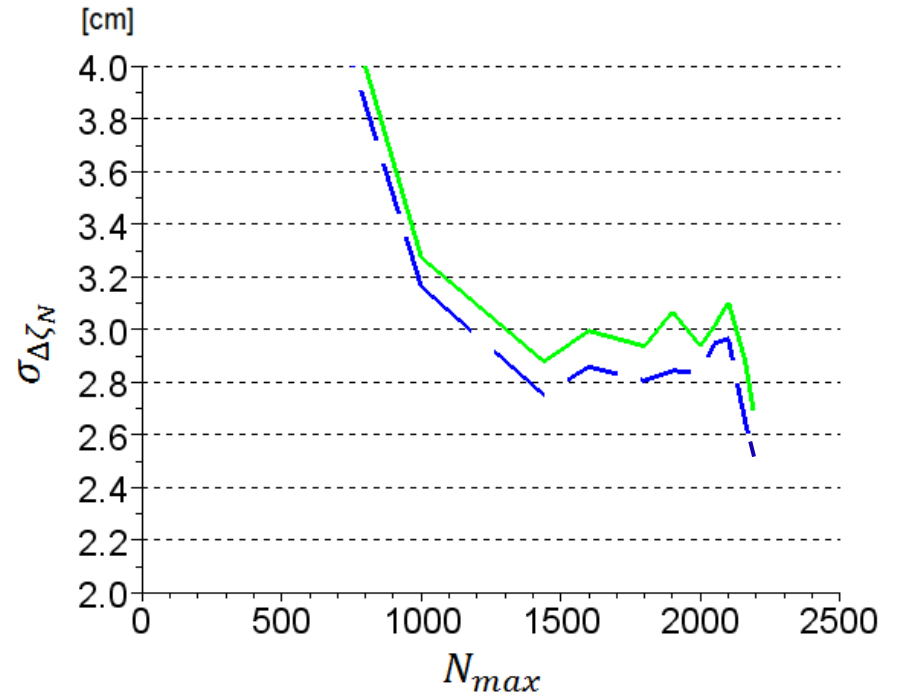
$$\Delta\zeta = \zeta_{GNSS/lev} - \zeta_{GM}$$

$$\Delta\delta g = \delta g_{MEASURED} - \delta g_{GM}$$

VALUES OF THE STANDARD DEVIATION ($\sigma_{\Delta\zeta_N}$) AS FUNCTION OF N_{max} VALUES

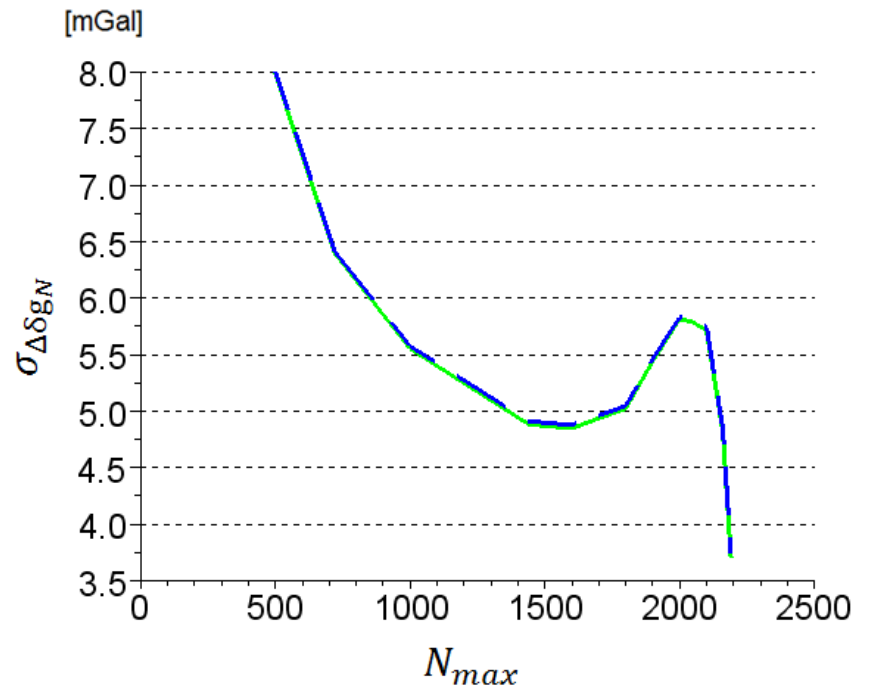
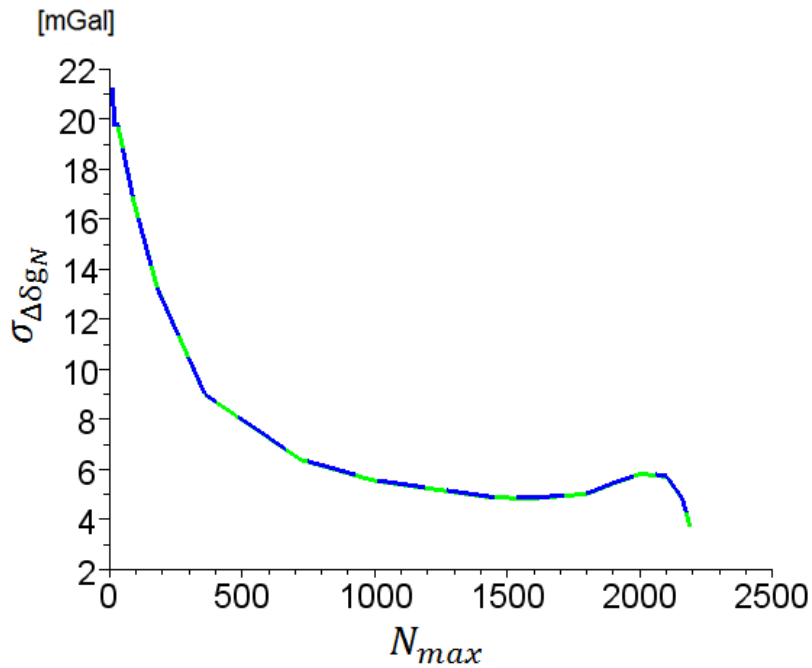


- EGM2008
- - - EIGEN-6C4



$$\Delta\zeta_N = \zeta_{GNSS/lev} - \zeta_{Nmax}$$

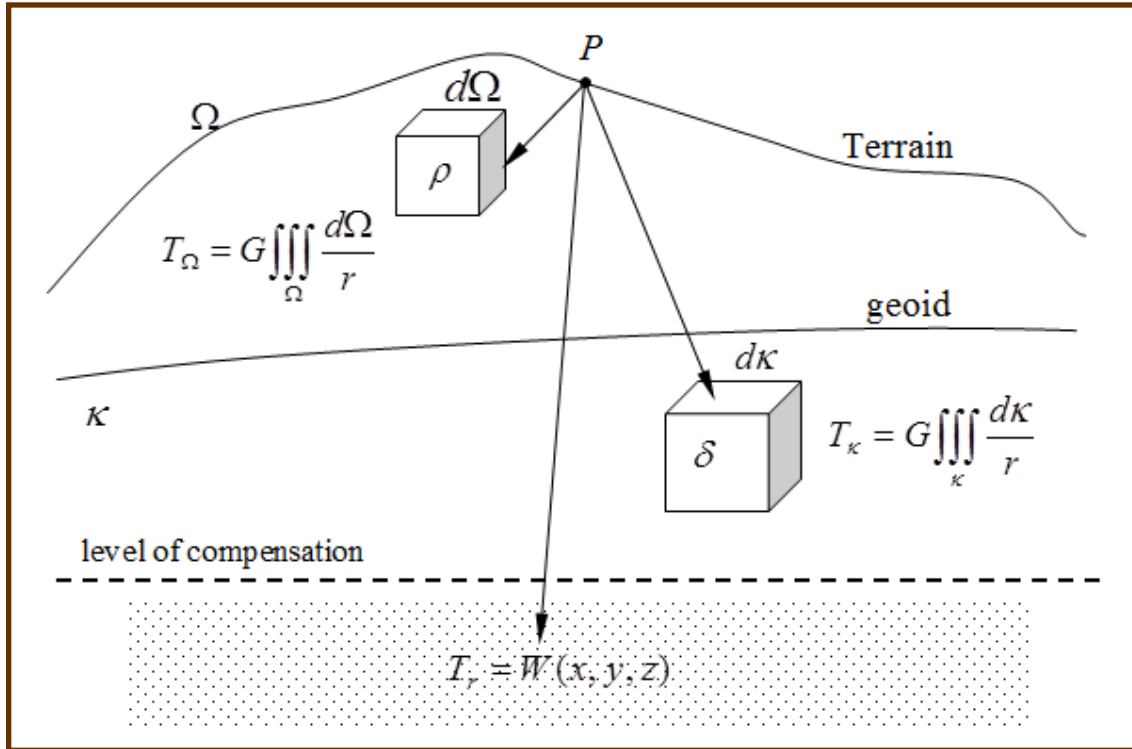
VALUES OF THE STANDARD DEVIATION ($\sigma_{\Delta\delta g_N}$) AS FUNCTION OF N_{max} VALUES



- EGM2008
- - - EIGEN-6C4

$$\Delta\delta g_N = \delta g_{IGSN71} - \delta g_{Nmax}$$

THE GGI METHOD OF LOCAL QUASIGEOID MODELLING



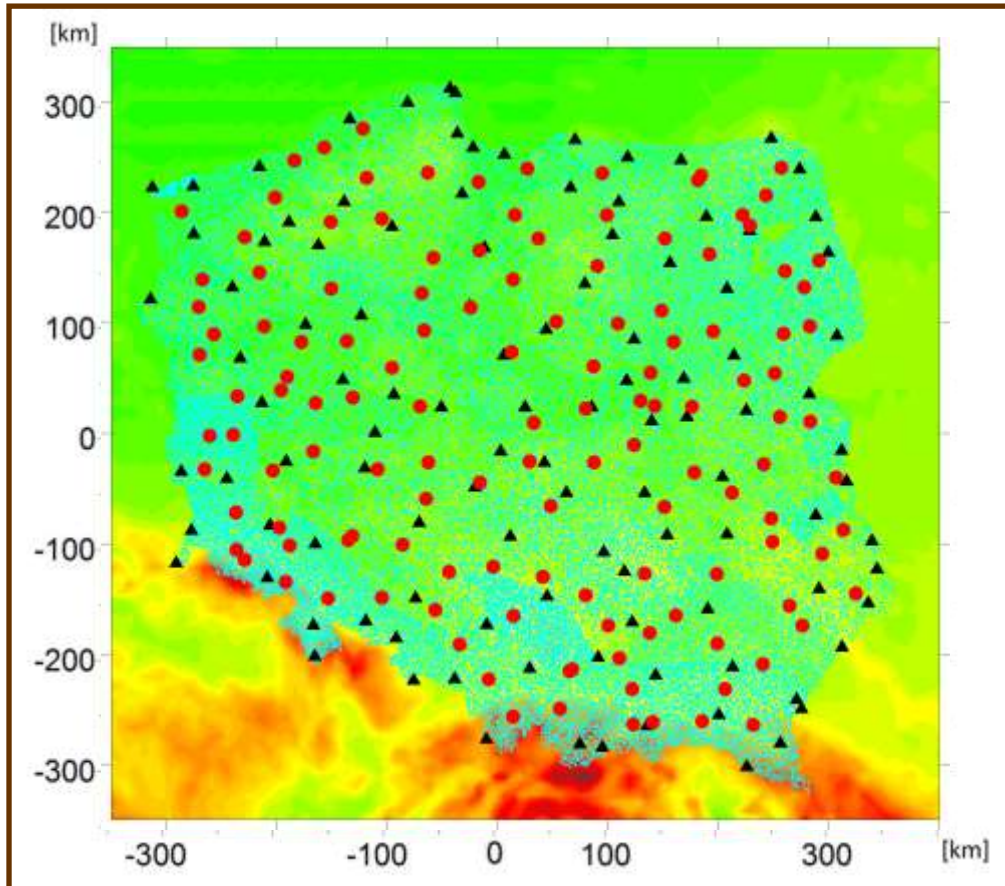
$$T_P = T_r + T_\Omega + T_\kappa$$

GGI method - local quasigeoid modelling method which uses Geophysical Gravity data Inversion technique

THE INPUT DATA

- Digital Terrain Model
- The Moho depth model
- Gravity data
- GNSS/levelling data

DATA USED FOR QUASIGEOID MODELLING BY GGI METHOD



Gravity points

▲ GNSS/Levelling data points

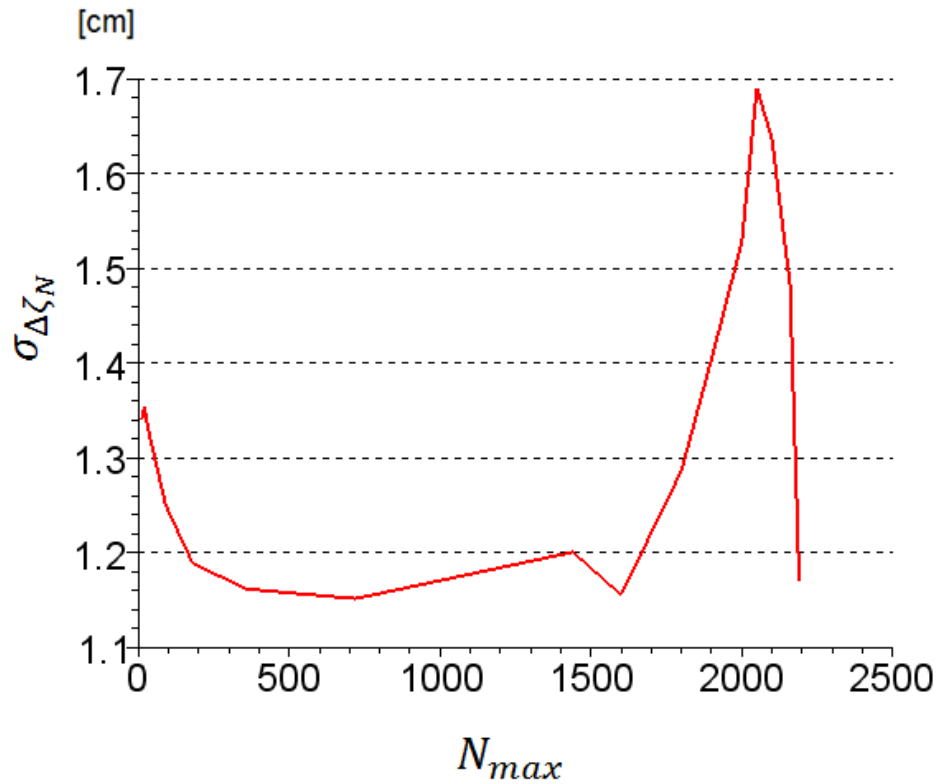
● GNSS/Levelling test points

BASIC ACCURACY PARAMETERS OF QUASIGEOID MODELS CALCULATED BY GGI METHOD
WITH THE USE EGM2008 AND EIGEN-6C4 GLOBAL MODELS

	EGM2008	EIGEN-6C4
Max($\Delta\zeta$) [cm]	3.9	3.9
Min($\Delta\zeta$) [cm]	-3.0	-2.9
stdev($\Delta\zeta$) [cm]	1.17	1.17

$$\Delta\zeta = \zeta_{GNSS/lev} - \zeta_{GGI}$$

VALUES OF THE STANDARD DEVIATION OF DIFFERATIONS $\Delta\zeta = \zeta_{GNSS/lev} - \zeta_{GGI}$
THE ζ_{GGI} VALUES CALCULATED WITH THE USE OF TRUNCATED EGM2008

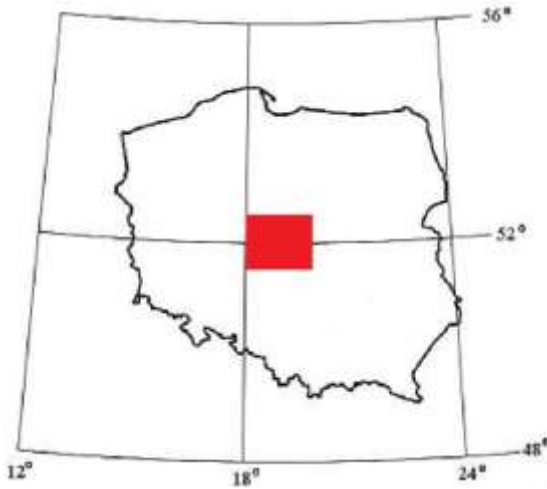


$$\Delta\zeta = \zeta_{GNSS/lev} - \zeta_{GGI}$$

DETAILED COURSE OF THE ANALYSED GLOBAL MODELS

For a small part of the central Poland ($51.5^\circ \leq \varphi \leq 52.5^\circ, 18^\circ \leq \lambda \leq 20^\circ$), dense grid ($\Delta\varphi = 0.01^\circ, \Delta\lambda = 0.02^\circ$) of the values $\zeta_{N_{max}}$ and $\delta g_{N_{max}}$ were determined. Because the changes in height anomalies ($\zeta_{N_{max}}$) were very small in relation to their values, a linear trend: $t_{N_{max}}(\varphi, \lambda) = a_0 + a_1\varphi + a_2\lambda$ for each set of $\zeta_{N_{max}}$ were estimated by least square method.

Subsequently residual height anomalies $\delta\zeta_{N_{max}} = \zeta_{N_{max}} - t_{N_{max}}$ were determined.



Dense grid of values $\zeta_{N_{max}}, \delta g_{N_{max}}$

Range of the grid: $51.5^\circ \leq \varphi \leq 52.5^\circ, 18^\circ \leq \lambda \leq 20^\circ$

Grid resolution: $\Delta\varphi = 0.01^\circ, \Delta\lambda = 0.02^\circ$

Truncation: $N_{max} \in \{1000, 1600, 2000, 2100, 2190\}$

Linear trend removed: $\delta\zeta_{N_{max}} = \zeta_{N_{max}} - t_{N_{max}}$

$$t_{N_{max}}(\varphi, \lambda) = a_0 + a_1\varphi + a_2\lambda$$

Contour maps of gravity disturbances $\delta g_{N_{max}}$ and residuals height anomalies $\delta \zeta_{N_{max}}$ determined for different N_{max} values from EGM08 (top) and EIGEN-6C4 (bottom) for the central part of Poland

$N_{max} = 1000$

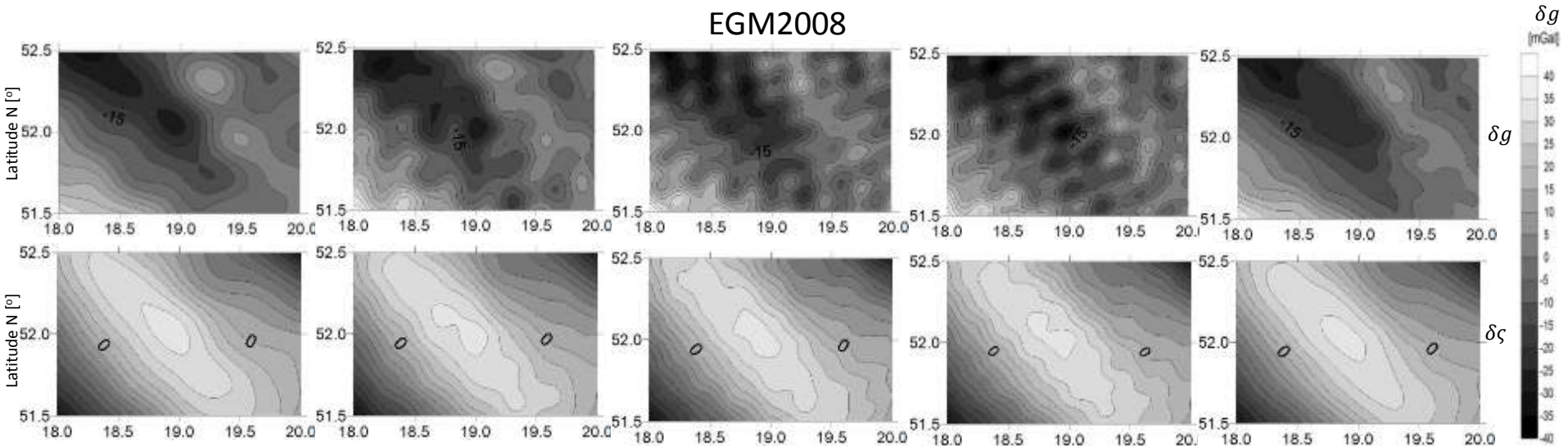
$N_{max} = 1600$

$N_{max} = 2000$

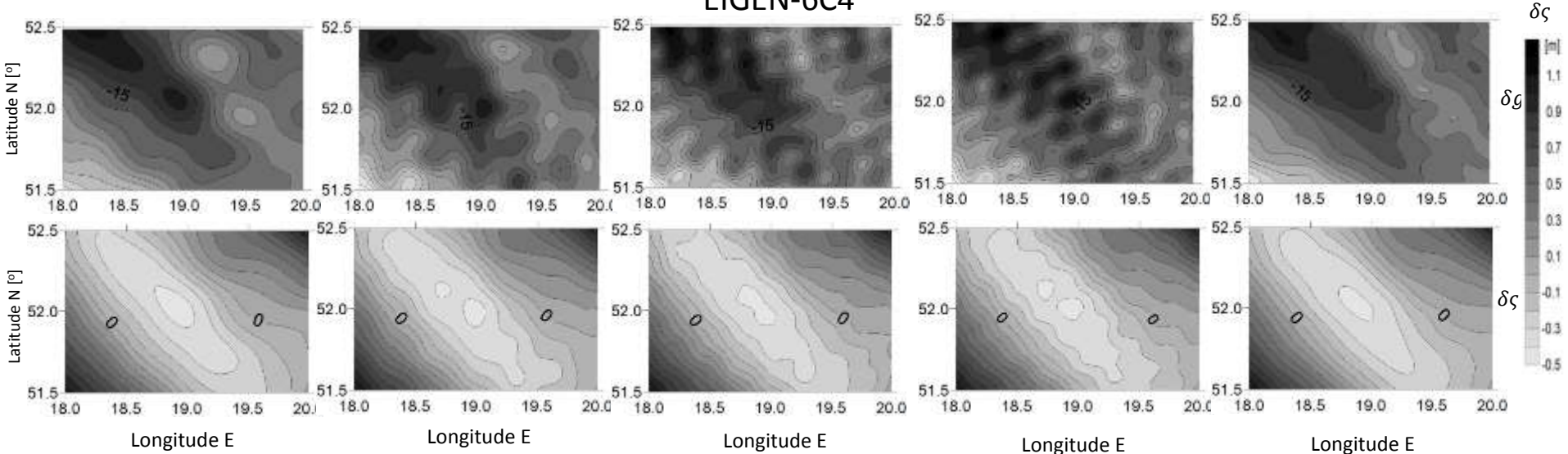
$N_{max} = 2100$

$N_{max} = 2190$

EGM2008



EIGEN-6C4



CONCLUSIONS

- Both tested models are very similar.
- The basic accuracy parameters of gravity disturbances are the same for both models ($\sigma_{\Delta\delta g} = \pm 3.7$ mGal).
- The basic accuracy parameters of height anomalies ($\sigma_{\Delta\zeta}$) are respectively:
 - ± 2.7 cm for EGM2008
 - ± 2.5 cm for EIGEN-6C4.

EIGEN-6C4 model is slightly more accurate (at the level of 8%)

- It was also noted an „unexpected behaviour” of both models when truncated models were studied. As a result of this, one can see a significant decrease in the accuracy of gravity disturbances and height anomalies of both models in the range of N_{max} about 2000-2100.

This indicates the need of using in geoid or quasigeoid modelling all coefficients of both models.



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Thank you for your attention